

# Improved turbidity estimation from local meteorological data for solar resourcing and forecasting applications

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## ABSTRACT

This work presents a new method to estimate atmospheric turbidity with improved accuracy in estimating clear-sky irradiance. The turbidity is estimated by machine learning algorithms using commonly measured meteorological data including ambient air temperature, relative humidity, wind speed and atmospheric pressure. The estimated turbidity is then served as the Linke Turbidity input to the Ineichen-Perez clear-sky model to estimate clear-sky global horizontal irradiance (GHI) and direct normal irradiance (DNI). When compared with the original Ineichen-Perez model which uses interpolated turbidity from the monthly climatological means, our turbidity estimation better captures its daily, seasonal, and annual variations. When using the improved turbidity estimation in the Ineichen-Perez model, the root mean square error (RMSE) of clear-sky GHI is reduced from  $24.02 \text{ W m}^{-2}$  to  $9.94 \text{ W m}^{-2}$ . The RMSE of clear-sky DNI is decreased from  $76.40 \text{ W m}^{-2}$  to  $29.96 \text{ W m}^{-2}$ . The presented method is also capable to estimate turbidity in partially cloudy days with improved accuracy, evidenced by that the corresponding estimated clear-sky irradiance has smaller deviation from measured irradiance in the cloudless time instants. In sum, the proposed method brings new insights about turbidity estimation in both clear and partially cloudy days, providing support to solar resourcing and forecasting.

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## 1. Introduction

Solar radiation reaching the Earth surface is either absorbed or scattered by the atmosphere based on the types and concentrations of the participating constituents and their radiative optical properties [1]. For solar energy conversion systems such as photovoltaics (PV) and concentrating solar power (CSP), ground level irradiance assessment and forecasting are crucial for their design and operation [2–5]. The attenuation of ground level solar irradiance is mainly caused by clouds, aerosols, water vapor, carbon dioxide and ozone [6], where clouds are the major modulator followed by aerosols and water vapor. However, the high temporal and spatial variations of the three major modulators as well as sensing difficulties of their concentrations [7] have posed considerable challenges for solar resourcing and forecasting applications. Therefore, a variety of clear-sky models have been developed over the years to estimate time varying ground level global horizontal

irradiance (GHI), direct normal irradiance (DNI) and diffuse horizontal irradiance (DHI) if there were no clouds in the sky. The clear-sky models have been used extensively to quantify the effects of local aerosols and water vapor, as well as to facilitate cloud identification and analysis for forecasting applications [1,8–10].

As summarized in Refs. [10,11], clear-sky models with different complexity and performance can be broadly classified into two main groups: physical models and empirical models. Physical models apply radiative transfer models (RTMs) to estimate the irradiance attenuation effect of atmospheric constituents, and the ground level solar irradiance can be obtained through integration of the attenuation caused by different atmospheric components [10]. Empirical models are based on simplified parameterizations of the attenuation processes [10], which estimate the clear-sky irradiance using some atmospheric parameters, such as the aerosol optical depth (AOD) and precipitable water in simplified Solis model [12], and the Linke turbidity ( $T_L$ ) in Ineichen-Perez model [13].

Physical models perform detailed analysis of the atmospheric attenuation processes, which generally lead to higher accuracy [11,14]. However, they require many inputs about local atmospheric conditions, some of them are not widely available. For instance, the

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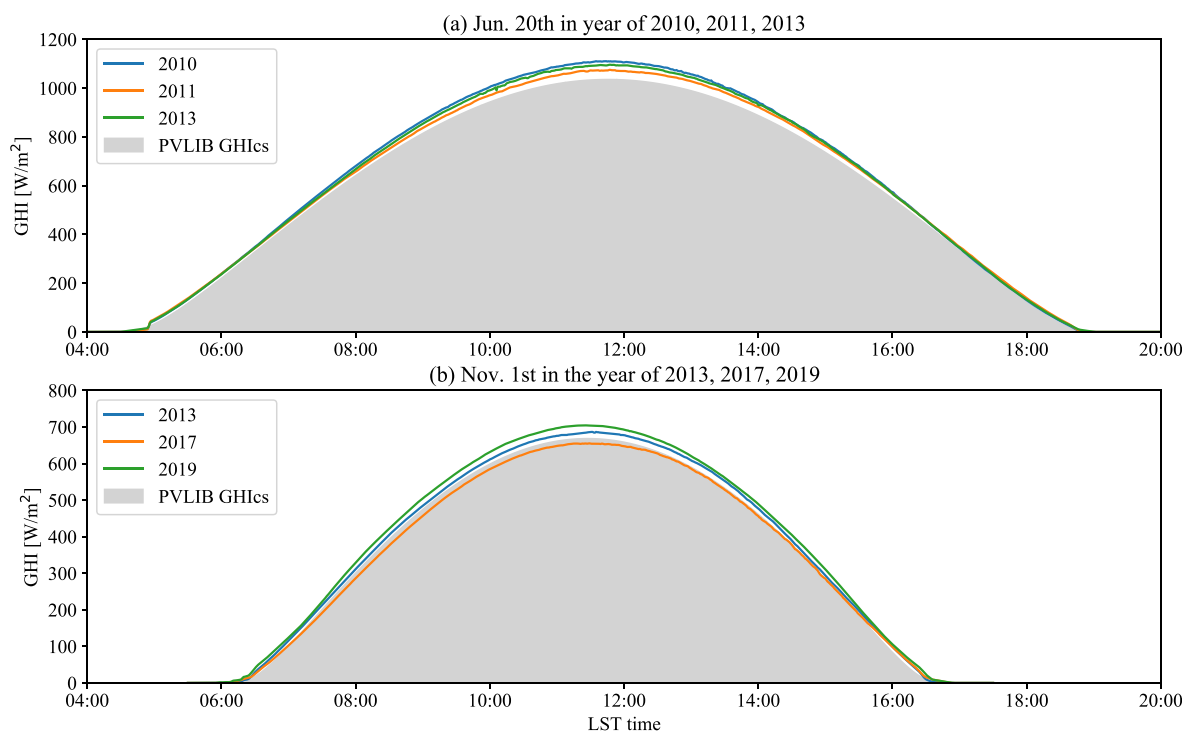
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REST2 model [15] has been verified as one of the most accurate clear-sky models [11,14], but the required information about the atmospheric constituents, such as AOD at 550 nm, column amount of ozone, nitrogen dioxide and precipitable water are difficult to obtain for most locations [14,16]. Yang [14] discussed the choice of clear-sky models in solar forecasting applications from the perspectives of accessibility, forecast performance and statistical properties. It is found that high-fidelity physical models like REST2 are not frequently used for solar forecasting due to its complexity, and no evidence suggests that physical models can lead to more accurate forecast results when compared with empirical models [14].

As a member of the empirical model family, the Ineichen-Perez model is extensively used in solar forecasting due to its simplicity [14]. The main input of the Ineichen-Perez model is the  $T_L$  factor, which is defined as the number of clean and dry atmospheres that produce the same attenuation equivalent to the real atmosphere [17]. The  $T_L$  factor quantifies the attenuation of aerosols and water vapor [18], which typically varies between 1 and 10 [10]. The  $T_L$  factor is available worldwide as monthly climatology value from the SoDa database [19]. In PVLIB [20], linear interpolations of the monthly values are applied to build daily  $T_L$  time series for each location when using the Ineichen-Perez clear-sky model.

The  $T_L$  factor is also directly used in other clear-sky models [10,21]. However, the invariant  $T_L$  factor based on the monthly climatology value and its linear interpolation cannot account for the short-term [22] and long-term variations [23] of atmospheric aerosols and water vapor concentrations, resulting in unsatisfying estimation of clear-sky irradiance [24]. The discrepancy of clear-sky irradiance obtained from  $T_L$  based clear-sky models and from measurements are observed in the studies by Moldovan et al. [21] and Polo et al. [22], and also noticeable when comparing the clear-sky solar irradiance measurements from Dessert Rock, Nevada (DRA) with PVLIB clear-sky model output (see Fig. 1).

Therefore, some studies were conducted to estimate  $T_L$  factor by different means with the aim to investigate its variations or to improve the estimation accuracy. Chaàbane et al. [7] adopted pyr heliometric measurements for the calculation of  $T_L$  factor in Tunisia during three summer months, where diurnal and monthly variations of  $T_L$  factor are observed. Polo et al. [22] estimated the daily  $T_L$  factor for clear days by using global irradiance measurements at solar noon and monthly mean  $T_L$  values. Using the estimated  $T_L$  to recalculate clear-sky solar irradiance results in a reduced root mean squared deviation (RMSD) when compared with using monthly mean values. The relative RMSD (rRMSD) decreases from 17.1% to 14.2% for the dataset of Baseline Surface Radiation Network (BSRN). For the dataset from Spanish Meteorological Agency (AEMet), the rRMSD reduces from 24.4% to 16.8%. Hove and Manyumbu [25] calculated the  $T_L$  factor based on daily GHIs and ESRA clear-sky model [26], which typically has a lower value than the monthly mean. Inman et al. [27] reported a method for daily average  $T_L$  estimation using broadband DNI measurements under cloudless skies, and then applied the estimated  $T_L$  in DNI forecasting during cloud-free periods under the assumption of a persistence of daily averaged  $T_L$  within the forecasting horizon. The relative RMSE (rRMSE) and relative mean bias error (rMBE) are smaller than 5% for both historical and forecasted values, which are much smaller than the error range (10–20%) of SoDa monthly means. Behar et al. [28] used ambient temperature and relative humidity to estimate  $T_L$  and solar irradiance via the estimated optical thickness of clean-dry atmosphere, water vapor and aerosol. The  $T_L$  estimation has a rRMSE of 10.22% and a rMBE of 1.31%, the rRMSE and rMBE of corresponding DNI estimate are 5.21% and 0.91%, respectively. Moldovan et al. [21] applied time dependent interpolation polynomials instead of a constant daily  $T_L$  factor to improve the clear-sky model. Two different interpolation polynomials are obtained for the  $T_L$  factor in warm and cold seasons, respectively.



**Fig. 1.** Comparison of measured clear-sky GHI (GHIs) with PVLIB GHIs of the same day in different years. PVLIB uses the  $T_L$  factor from its look-up table, which is based on constant monthly climatology value. The PVLIB GHIs remains the same on the same day of different years, while the measured GHIs are not, indicating  $T_L$  factor also varies on a long term (i.e., yearly) basis.

The result shows that the relative error is reduced from 8.12% to 4% in the warm season and from 5.02% to 4.15% in the cold season.

The derivation of  $T_L$  based on irradiance and meteorological measurements offers a simpler way to estimate  $T_L$  without the detailed information about aerosol and water vapor contents. However, the methods summarized above still have some shortcomings to overcome. For example, the clear-sky irradiance measurements are not available in a cloudy day, and only using the GHIs at the solar noon presented by Polo et al. [22] may lead to errors in estimating the clear-sky irradiance in other periods such as solar mornings, evenings and cloudy days. The method also cannot be used for locations without irradiance measurements. For the method presented by Behar et al. [28], using ambient temperature and relative humidity to estimate perceptible water and AOD may result in error accumulations in estimating  $T_L$ . For the study by Modlovan et al. [21], the  $T_L$  interpolation polynomials for warm and cold seasons are not capable of accounting for year-to-year  $T_L$  variation as shown in Fig. 1.

Therefore, we propose a new  $T_L$  estimation method to estimate high-fidelity  $T_L$  with the consideration of its short-term and long-term variations, and without the data dependence on local real-time irradiance measurements. The  $T_L$  factor is proposed to be estimated using local meteorological data by machine learning (ML) algorithms. In the following sections, data processing and proposed methodology are presented in Section 2. Section 3 presents the results and discussions of  $T_L$  and corresponding clear-sky irradiance estimations. The key findings and recommendations are summarized in Section 4.

## 2. Methodology of turbidity estimation

This section presents the data and methods used for  $T_L$  derivation and estimation. The  $T_L$  derivation is performed by applying Ineichen-Perez clear-sky model (PVLIB) reversely, i.e., taking the 1-min averaged GHIs as the input to compute the ‘ground truth’  $T_L$ . Then the derived minute-wise  $T_L$  time series is further averaged on the basis of daily, hourly and 5-min as the ML model training targets. The input meteorological data is also averaged with the same time basis for model training, tuning and testing. Finally, the trained model is applied to estimate the  $T_L$  for GHIs estimation. The flowchart of the method for estimating the  $T_L$  and clear-sky irradiance is shown in Fig. 2. The  $T_L$  derivation and estimation can

also be applied to clear-sky DNI (DNIs), which will be discussed in Section 3.3.

### 2.1. Data selection

The data used in this work is from DRA, one of the Surface Radiation Budget Network (SURFRAD) stations [29]. DRA has a latitude of 36.62373°N, a longitude of 116.01947°W, an elevation of 1007 m, and a time zone of UTC-8 (8 h difference than coordinated universal time (UTC)). High resolution solar irradiance and meteorological data collected from year 2000–2020 are used in this work. Among the diverse variables in the comprehensive dataset, measurements of the downwelling global solar irradiance (GHI), direct normal irradiance (DNI), ambient air temperature ( $T_a$ ), relative humidity ( $\phi$ ), wind speed ( $V$ ) and local atmospheric pressure ( $P_a$ ) are selected to build and test the proposed  $T_L$  estimation model. The selected data has high temporal resolutions (3-min averaged from year 2001–2008, and 1-min averaged from year 2009–2020) and its quality is carefully controlled.

DRA is chosen among the seven SURFRAD stations due to its high occurrence of cloudless days, which could provide adequate learning samples for the development and validation of the proposed  $T_L$  estimation model. The same methodology can be applied to other locations if sufficient data is given.

### 2.2. Selection of clear-sky days

The clear-sky irradiance is defined as the incident radiation at the Earth’s surface under the conditions that would occur under a perfect ‘‘cloudless sky’’ [30]. The presence of clouds in the sky, especially when clouds obscure the Sun disc, will greatly affect the surface solar radiation, resulting in irradiance fluctuations. Since  $T_L$  is a factor that quantifies the attenuation of solar irradiance by atmospheric constituents (especially water vapor and aerosols) under cloud-free conditions, we only select clear-sky days for model development and validation.

The clear-sky days are selected following the approach developed by Long and his collaborators [31–33], and the clear-sky labels provided by RadFlux algorithm [31,33] are publicly available on the website of SURFRAD network. Specifically, the days will be labeled as ‘‘clear-sky day’’ if most of the time instants within the day are ‘‘clear’’ as detected by solar shortwave irradiance measurements ( $\lambda < 4 \mu\text{m}$ ), or detected by atmospheric longwave irradiance

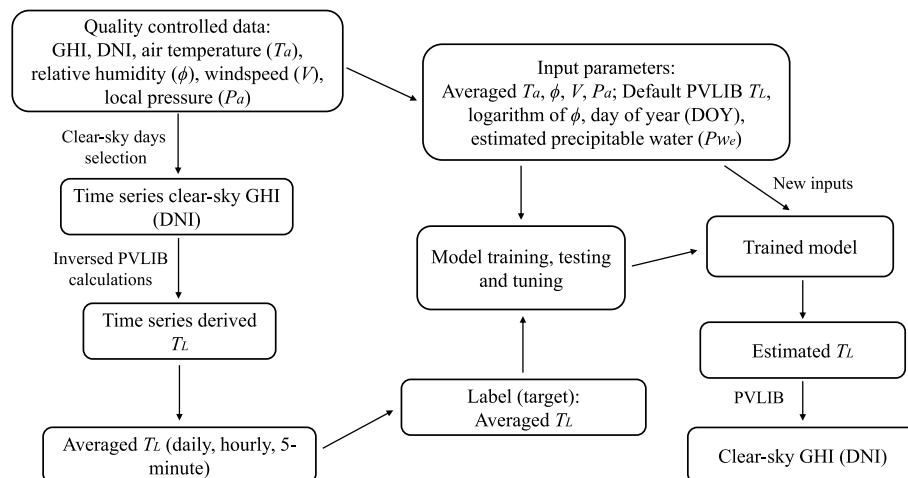


Fig. 2. An overview of the method to derive and estimate  $T_L$ . The model for estimating the daily, hourly and 5-min  $T_L$  is trained independently, and the input meteorological data is also averaged on the same time basis.

measurements ( $\lambda > 4 \mu\text{m}$ ). The shortwave clear-sky detection algorithm has a  $160^\circ$  field of view, so the clear-sky instants with high solar zenith angles could not be detected [32]. During the daytime, the presence of clouds is more noticeable in the shortwave spectrum when compared with the longwave spectrum [33]. Therefore, most of the clear-sky instants detected by the longwave RadFlux algorithm [33] are in the nighttime. To verify the RadFlux clear-sky labels, we performed an additional manual check by comparing the measured GHI, clear-sky GHI estimated by RadFlux [34] and Ineichen–Perez clear-sky model from PVLIB [20]. Then, clear-sky detection for periods with high solar zenith angles are performed, and wrongly labeled clear-sky days are removed, as demonstrated in Fig. 3. Our training and testing datasets only contain the clear-sky days that pass both the RadFlux and manual checks (examples are presented in Fig. 3).

### 2.3. Derivation of ‘ground truth’ turbidity for model training

We adapt the methodology documented in PVLIB [20] to derive ‘ground truth’  $T_L$  factor for model development. At each clear-sky time instance, the ‘ground truth’  $T_L$  factor is derived from measured GHICs values by inverting the following equation provided in PVLIB (proposed in Ref. [13]),

$$\text{GHI}_{\text{cs}} = c_1 \cdot I_0 \cdot \cos(\theta) \cdot \exp(-c_2 \cdot \text{AM} \cdot (f_1 + f_2 \cdot (T_L - 1)))$$

Then the derived  $T_L$  based on GHICs measurement is,

$$T_L = \left[ \ln\left(\frac{\text{GHI}_{\text{cs}}}{c_1 \cdot I_0 \cdot \cos(\theta)}\right) / (-c_2 \cdot \text{AM}) - f_1 \right] / f_2 + 1 \quad (1)$$

$T_L$  could also be derived from DNICs by inverting the following equations from PVLIB,

$$B_1 = I_0 \cdot b \cdot \exp(-0.09 \cdot \text{AM} \cdot (T_L - 1))$$

$$B_2 = \text{GHI}_{\text{cs}} \cdot \left[ \left( 1 - \frac{0.1 - 0.2 \cdot \exp(-T_L)}{0.1 + 0.882/f_1} \right) / \cos(\theta) \right]$$

$$\text{DNI}_{\text{cs}} = \text{Minimum}(B_1, B_2)$$

Then the derived  $T_L$  based on DNICs measurements is,

$$T_L = \ln\left(\frac{\text{DNI}_{\text{cs}}}{I_0 \cdot b}\right) / (-0.09 \cdot \text{AM}) + 1, \text{ (when } B_1 < B_2) \quad (2)$$

$$T_L = -\ln\left[\left(0.1 - \left(1 - \frac{\text{DNI}_{\text{cs}} \cdot \cos(\theta)}{\text{GHI}_{\text{cs}}}\right) \cdot (0.1 + 0.882/f_1)\right) / 0.2\right], \text{ (when } B_1 > B_2) \quad (3)$$

with:

$$\text{AM} = \left( \frac{1}{\cos(\theta) + 0.50572 \cdot (6.07995 + (90 - \theta)^{-1.6364})} \right) \cdot \frac{P_a}{101325}$$

$$c_1 = 5.09 \cdot 10^{-5} \cdot h + 0.868$$

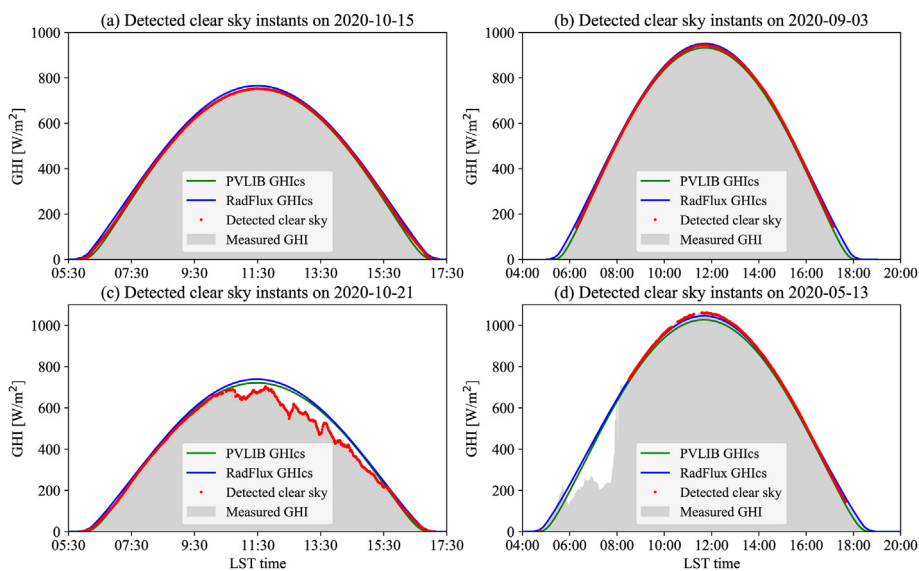
$$c_2 = 3.92 \cdot 10^{-5} \cdot h + 0.0387$$

$$f_1 = \exp(-h / 8000)$$

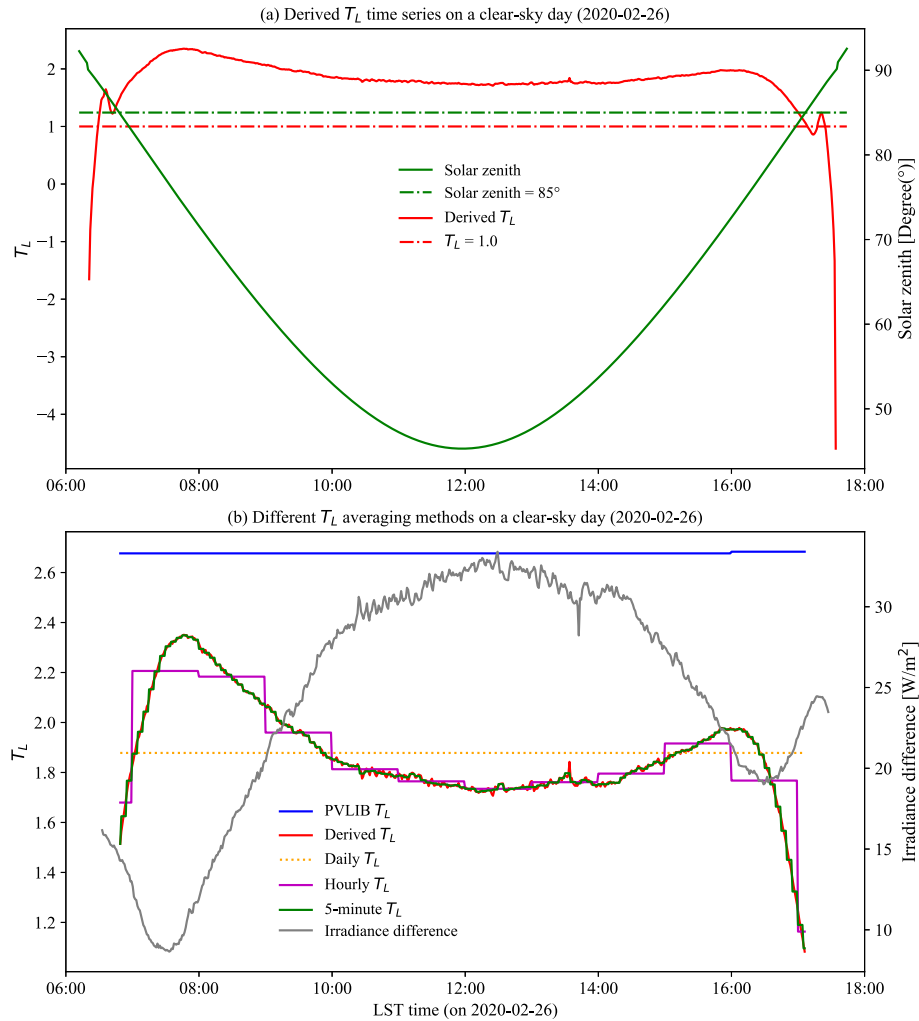
$$f_2 = \exp(-h / 1250)$$

$$b = 0.664 + 0.163/f_1$$

where  $\text{GHI}_{\text{cs}}$  [ $\text{W m}^{-2}$ ] is the measured clear-sky GHI.  $\text{DNI}_{\text{cs}}$  [ $\text{W m}^{-2}$ ] is the measured clear-sky DNI.  $B$  [ $\text{W m}^{-2}$ ] is the normal beam clear-sky radiation.  $c_1$ ,  $c_2$ ,  $f_1$ ,  $f_2$ ,  $b$  are altitude-dependant coefficients,  $I_0$  [ $\text{W m}^{-2}$ ] is the solar constant,  $\theta$  [ $^\circ$ ] represents the solar zenith angle,  $\text{AM}$  is the absolute airmass,  $T_L$  is the Linke Turbidity factor,  $P_a$  [Pa] is the local atmospheric pressure, and  $h$  [m] is local altitude.



**Fig. 3.** Examples of clear-sky days selection in DRA. Measured GHI data is from SURFRAD, GHICs are computed by both RadFlux and PVLIB. The clear-sky labels are from RadFlux. (a) A detected full clear-sky day. (b) A detected clear-sky day with high solar zenith periods not labeled. (c) A wrongly labeled clear-sky day which is removed by manual check. (d) A typical partly cloudy day.



**Fig. 4.** Derived  $T_L$  time series with different averaging modes for a clear-sky day with respect to local standard time (LST). (a) Derived  $T_L$  time series on 2020-02-26. The derived  $T_L$  shows high variations and unrealistic values in the periods with high solar zenith angles. (b) Averaged  $T_L$  on different time basis and the irradiance difference between measured GHIs and PVLIB GHIs during the day.

Fig. 4 (a) illustrates the GHIs-derived  $T_L$  in a randomly selected clear-sky day (one can find similar results in any other clear-sky days). Unlike the  $T_L$  factor used in PVLIB default calculations, the derived  $T_L$  factor is not a constant but varies during the day. Note that for periods with large solar zenith angle (greater than  $85^\circ$ ), the derived  $T_L$  has large variations and unrealistic values (less than 1.0 and even negative), which are resulted from the applicable limitations of Eq. (1). Therefore, in the following sections when derived  $T_L$  time series are temporally averaged, the instances from the period when the solar zenith is greater than  $85^\circ$  are not included. Fig. 4 (b) shows the  $T_L$  time series averaged on different time basis, where the daily averaged  $T_L$  is much lower than the value used in PVLIB. The asymmetry of estimated  $T_L$  with respect to the zenith angle is observed in Fig. 4 (b), especially during morning and evening periods. This is possibly due to the high airmass effect, where the small difference in measured clear-sky irradiance will result in large discrepancy in the derived  $T_L$  values. In addition, the profile of clear-sky irradiance is not perfectly symmetric as shown in Fig. 4 (b) that the measured GHIs in the morning (e.g., 6:00–8:00) is smaller than the ones near the evening (e.g., 16:00–18:00). Meanwhile, the water vapor content in the atmosphere is usually higher in the morning, which results in higher  $T_L$  values and thus lower GHIs.

The averaged  $T_L$  derivations are then used to recalculate the 1-min averaged GHIs using PVLIB, which shows noticeable improvement in estimating GHIs, as shown in Fig. 5. In the clear-sky days of 2019 and 2020 (a total of 84 days are identified as clear), using PVLIB  $T_L$  generally underestimates the GHIs with a mean bias error (MBE) of  $-20.48 \text{ W m}^{-2}$  and a root mean square error (RMSE) of  $24.02 \text{ W m}^{-2}$ , when computed using 1-min averaged data when solar zenith angle is smaller than  $85^\circ$ . Using derived daily mean  $T_L$  yields a GHIs estimation with overall MBE of  $0.34 \text{ W m}^{-2}$  and RMSE of  $6.74 \text{ W m}^{-2}$ , a 98.3% reduction in MBE and 71.9% decrease in RMSE. As the time resolution increases, the recalculated results become better as expected. Hourly mean  $T_L$  produces a RMSE of  $2.81 \text{ W m}^{-2}$  and a MBE of  $0.05 \text{ W m}^{-2}$  for GHIs estimation. The 5-min averaged  $T_L$  gives an estimation of GHIs with the lowest MBE of  $0.01 \text{ W m}^{-2}$  and lowest RMSE of  $0.55 \text{ W m}^{-2}$ . In general, using daily mean  $T_L$  can successfully correct the bias in estimating GHIs and reduce RMSE by 71.9%. Using temporally finer hourly and 5-min averaged  $T_L$  can further reduce the RMSE in GHIs estimations, but they also substantially increased the size of training data in the following ML based  $T_L$  estimation models.

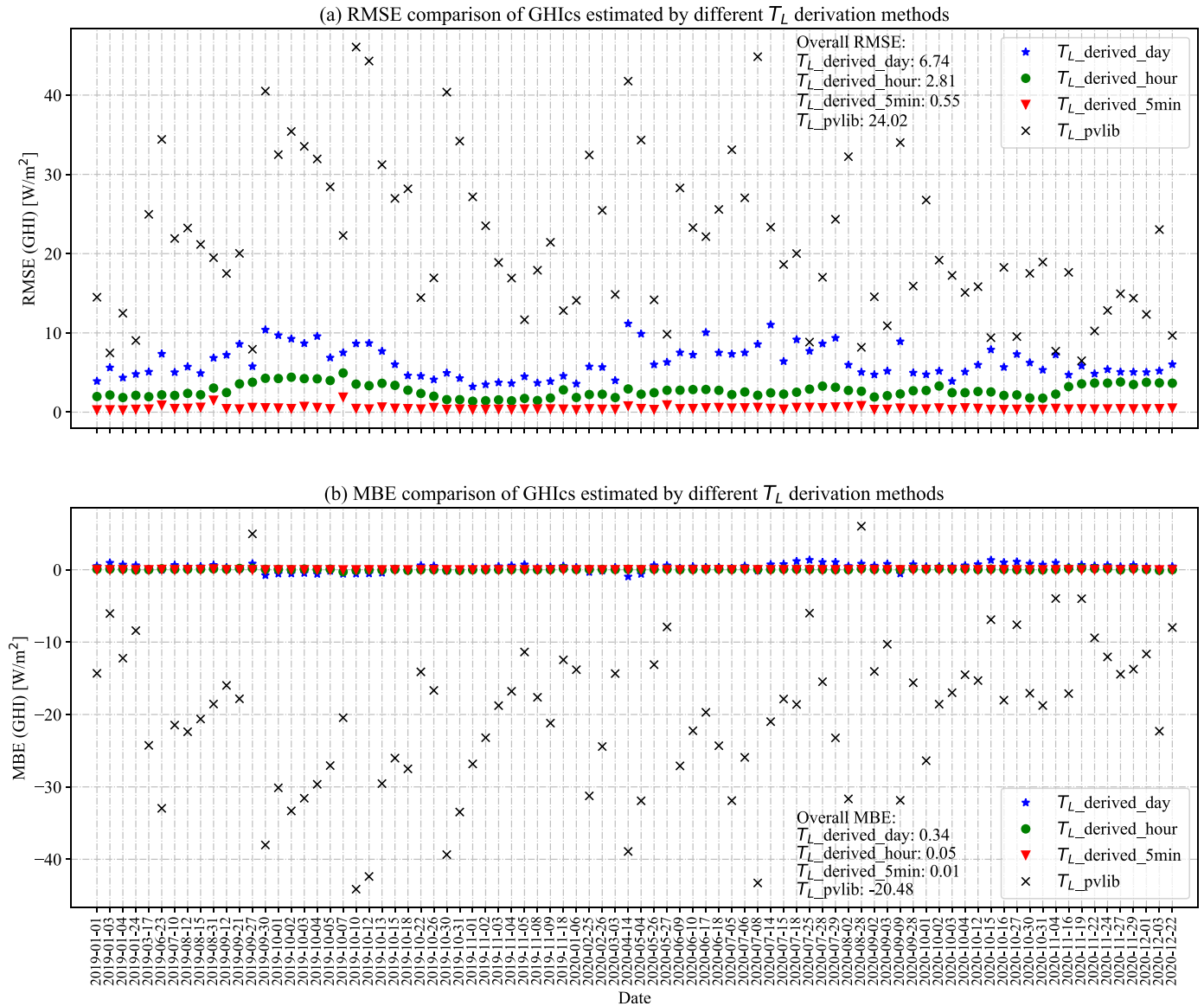


Fig. 5. Comparison of GHICs estimation using different  $T_L$  averaging modes for the clear-sky days in year 2019 and 2020. (a) Daily RMSE of 1-min averaged GHICs estimation. (b) Daily MBE of 1-min averaged GHICs estimation. All the derived  $T_L$  regardless of averaging mode produce more accurate GHICs than PVLIB.

#### 2.4. Turbidity estimation from local meteorological data

The previous section demonstrates that improving  $T_L$  estimations could substantially improve the accuracy of GHICs estimations. However, the  $T_L$  is derived from GHICs measurements, which is not known as a priori in real-time applications. Therefore, we propose to use ML methods with widely available meteorological measurements to estimate local  $T_L$ .

We use three independent ML models for daily, hourly and 5-min averaged  $T_L$  factors. The label (target) is the averaged  $T_L$  derivations from Section 2.3, and the input parameters are: the default PVLIB  $T_L$ , ambient air temperature  $T_a$ , relative humidity  $\phi$  (and its logarithm), wind speed  $V$ , atmospheric pressure  $P_a$ , day of year (DOY), and estimated precipitable water  $Pw_e$ . The meteorological time series (i.e., air temperature, relative humidity, wind speed, pressure) are averaged on the same time basis as the  $T_L$ . The PVLIB  $T_L$  is adapted to the corresponding time resolution as well. The logarithm of relative humidity is based on its averaged value, and the estimated precipitable water is calculated from the averaged

temperature and the averaged relative humidity using the empirical model proposed by Gueymard [35,36] with the following equations.

$$Pw_e = 0.1 \cdot H_v \cdot \rho_v$$

$$H_v = 0.4976 + \frac{1.5265 \cdot T_a}{273.15} + \exp\left(\frac{13.6897 \cdot T_a}{273.15} - 14.9188 \cdot \left(\frac{T_a}{273.15}\right)^3\right)$$

$$\rho_v = 216.7 \cdot \phi \cdot e_s / T_a$$

$$e_s = \exp\left(22.330 - 49.140 \cdot \frac{100}{T_a} - 10.922 \cdot \left(\frac{100}{T_a}\right)^2 - 0.39015 \frac{T_a}{100}\right)$$

where  $Pw_e$  [cm] is the estimated precipitable water.  $H_v$  [km] is the apparent water vapor scale height.  $\rho_v$  [ $g\ m^{-3}$ ] is the surface water vapor density.  $\phi$  [%] is the relative humidity.  $e_s$  [millibar] is the saturation water vapor pressure.  $T_a$  [ $^{\circ}C$ ] is the ambient air temperature.

ML technique is a powerful tool in regression modelling, which can model the relations between input features and target especially when the representation is complicated. ML algorithms have been widely used in classification, prediction and pattern recognition applications [37]. Here, we apply and compare Linear Regression (LR), Random Forest (RF), and Multilayer Perceptron (MLP) for  $T_L$  estimation, which are three commonly available and intensively used methods in real applications.

LR involves a linear combination of the input variables, which may have significant limitations for pattern recognition, particularly for problems with high dimensionality [37]. Therefore, linear model is extended by considering linear combinations of fixed nonlinear functions (basis function) of the input variables. Polynomial (powers of input variables) regression is one example of the extended linear models [37]. Although linear models are considered relatively simple and might not be suitable for high-dimensional problems, they have good analytical properties and form the fundamental for more advanced models [37]. Here we apply LR as a reference method in estimating  $T_L$ .

RF regressor is an ensemble method that combines several randomized regression decision trees to achieve a better performance [38]. RF is a bagging technique, all the involved decision trees are built in parallel and depend on the random vectors sampled from the training dataset. The predictions are averaged using bootstrap aggregation, which is one of the most computational-efficient methods to improve stability of the estimates [38]. RF models have been demonstrated to be robust predictors for both small sample sizes and data with high dimensionality [38].

MLP is also known as feed-forward neural network, which consists of an input layer, one or more hidden layers and one output layer [37]. MLP networks have high flexibility in approximation and can easily extend the structure by adding more hidden layers. MLP networks are trained and the parameters are obtained by back propagation [37]. There are different nonlinear activation functions of hidden layer(s), which could differ for different applications.

Data from 2000 to 2018 is used as the training set (20% of which

is for validation) and data from 2019 to 2020 is used for testing. The model hyperparameters are tuned by using tenfold cross-validation method. The error evaluation metrics are MBE, RMSE and their normalized counterparts. All the above-mentioned ML models are adapted from Scikit-learn [39] and PyCaret [40], where more details regarding the applied algorithms can be found.

### 3. Results and discussion

The best ML model is selected separately for daily, hourly and 5-min averaged  $T_L$  estimation, and the overall corresponding 1-min averaged GHICs estimation results for clear-sky days in year 2019 and 2020 are presented in Table 1. Compared with the GHICs recalculation, GHICs based on the estimated  $T_L$  yields slightly larger MBE and RMSE. Although 5-min averaged  $T_L$  has the best performance for GHICs recalculations, but the fine temporal resolution does not show much superior results in GHICs estimation. Estimating hourly averaged  $T_L$  results to better GHICs estimation with an MBE of  $1.45\ W\ m^{-2}$  and a RMSE of  $9.62\ W\ m^{-2}$ . Using daily averaged  $T_L$  achieves a comparable result with slightly larger MBE of  $2.09\ W\ m^{-2}$  and RMSE of  $9.94\ W\ m^{-2}$ . Given that less complexity and computational resource are required for using daily averaged  $T_L$ , the subsequent results and discussion are based on the daily averaged  $T_L$  and the associated model.

#### 3.1. Estimations of daily turbidity and 1-minute averaged GHICs in clear-sky days

When compared with the monthly climatology mean of  $T_L$ , the derived daily  $T_L$  generally has a lower value and has a much higher fluctuation (see Fig. 6 (a)). The  $T_L$  values of 2020 is different from the year of 2019, which indicates the  $T_L$  also has a yearly variation. This long-term fluctuation of  $T_L$  is possibly caused by pollution [23] and the dynamics of aerosols and water vapor in the atmosphere [22]. Among the applied ML algorithms, MLP regressor gives the best results as shown in Table 2 and Fig. 6, with comparatively lower testing RMSE and MBE values. The normalized RMSE (nRMSE) of  $T_L$  estimation from all the ML models are around 10%. The learning curve of MLP regressor is shown in Fig. 7.

Fig. 8 (a) presents the sensitivity analysis of the meteorological inputs for the MLP model. The estimated  $T_L$  increases when the temperature and relative humidity become higher, while the increases in wind speed and pressure lead to a drop in the  $T_L$  estimation. Wind speed is the least sensitive parameter, so its impact on the  $T_L$  estimation is limited. Relative humidity and temperature have comparatively larger influence than wind speed, and temperature is a more crucial input for  $T_L$  estimation compared with relative humidity. Regards to local pressure, it does not have large variance as shown in Fig. 8 (b), so either increase or decrease pressure by 10% would lead it to be out of its min-max range. Since the MLP model is trained based on data samples with a small range of pressure variation, the out-of-range pressure will produce unrealistic  $T_L$  estimation. This is why the pressure shows the relatively larger sensitivity. In practical applications, the pressure of a certain place has limited variation, so its influence on  $T_L$  estimation also remains limited.

When using the estimated daily  $T_L$  from the MLP model to estimate 1-min averaged GHICs, most of the tested clear-sky days in the year of 2019 and 2020 show noticeable improvements when compared with the PVLIB GHICs in terms of RMSE and MBE (see Fig. 9). The overall RMSE of GHICs estimation using the MLP-estimated  $T_L$  in 2019 and 2020 is  $9.94\ W\ m^{-2}$ , which is slightly higher than the RMSE ( $6.74\ W\ m^{-2}$ ) of GHICs recalculation from the derived  $T_L$ , but much lower than the RMSE of  $24.02\ W\ m^{-2}$  from PVLIB. Note that there are some cases of model underperformance,

**Table 1**

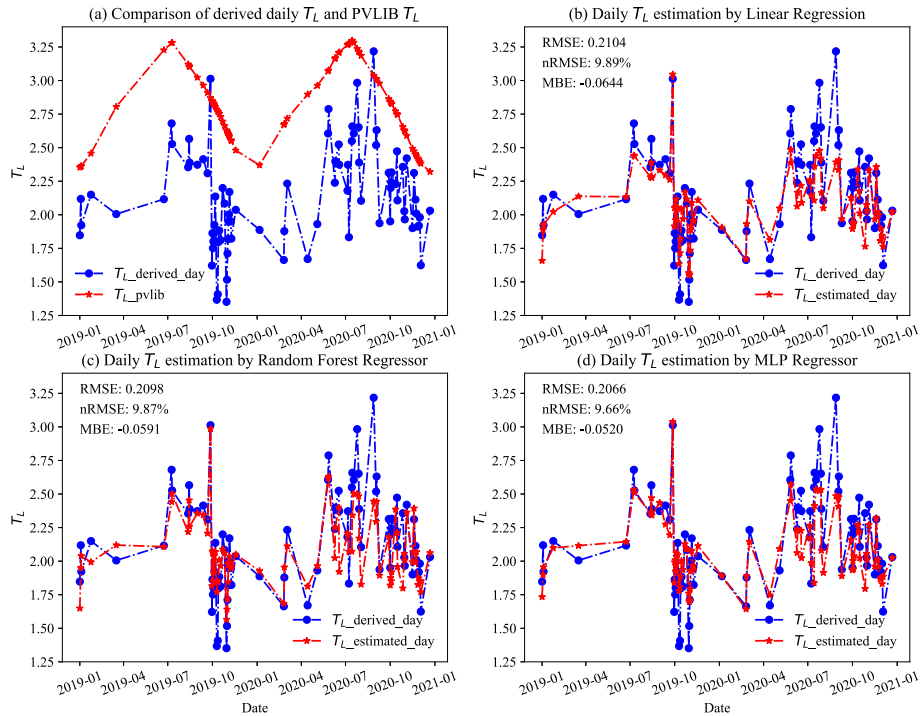
Comparison of 1-minute averaged GHICs recalculations and estimations using derived and estimated  $T_L$  for clear-sky days in the year of 2019 and 2020. PVLIB results are presented here for reference.

$T_L$	GHICs recalculations <sup>a</sup>		GHICs estimations <sup>b</sup>	
	RMSE [ $W\ m^{-2}$ ]	MBE [ $W\ m^{-2}$ ]	RMSE [ $W\ m^{-2}$ ]	MBE [ $W\ m^{-2}$ ]
Daily mean	6.74	0.34	9.94	2.09
Hourly mean	2.81	0.05	9.62	1.45
5-min mean	0.55	0.01	10.28	-0.01
PVLIB <sup>c</sup>	24.02	-20.48	24.02	-20.48

<sup>a</sup> GHICs recalculations are based on the averaged  $T_L$  factors derived from GHICs.

<sup>b</sup> GHICs estimations are based on the estimated  $T_L$  factors from the ML (MLP) models with meteorological parameters as input.

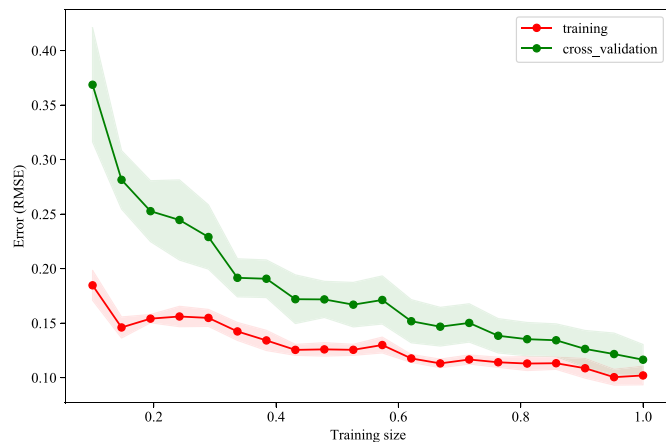
<sup>c</sup> PVLIB uses the daily interpolated  $T_L$  based on the monthly climatological  $T_L$  map.



**Fig. 6.** Comparison of the derived daily  $T_L$  and PVLIB  $T_L$  and the performance of applied ML methods. (a) The comparison of derived  $T_L$  and PVLIB  $T_L$ . The comparison between derived  $T_L$  and estimated  $T_L$  from different methods (b) Linear Regression (c) Random Forest Regressor and (d) MLP Regressor.

**Table 2**  
Training and testing errors of the applied ML algorithms for  $T_L$  estimation.

ML algorithm	Training		Testing	
	RMSE	MBE	RMSE	MBE
LR	0.245 2	0.000 0	0.210 4	-0.064 4
RF	0.212 7	-0.000 4	0.209 8	-0.059 1
MLP	0.233 9	0.004 4	0.206 6	-0.052 0



**Fig. 7.** Learning curve of MLP regressor with tenfold cross validation. The dots represent mean values, and the related shadows reflect the standard deviation.

which are likely due to that PVLIB  $T_L$  is already close to the derived  $T_L$ . Nevertheless, the GHICs estimation using the estimated  $T_L$  factor has an overall better performance compared with PVLIB, which uses unmodified  $T_L$  based on the monthly climatology values, especially when the PVLIB GHICs and measured GHICs have large discrepancy.

### 3.2. Estimations of daily turbidity and 1-minute averaged GHICs in partially clear days

Furthermore, we test our  $T_L$  estimation model in partially cloudy days when not all periods are cloudless throughout the day. As shown in Fig. 10, the model is also applicable to estimate  $T_L$  in this case and the corresponding 1-min averaged GHICs estimation shows better agreement when compared with PVLIB for the clear-sky instants during the day. The potential explanation to this phenomenon is that the presence of clouds in partially clear days has limited effect on local meteorological parameters as well as ground level aerosols and water vapor concentrations. Accordingly, using local meteorological measurements (e.g., temperature, relative humidity) to estimate GHI with the presence of clouds may not be effective. Note that the phenomenon could be different in fully overcast days as the meteorological parameters might be affected, which needs further investigation. Nevertheless, the trained model works for the partially cloudy days, which would provide more accurate clear-sky irradiance during those periods for solar resourcing and forecasting applications. In addition, since the ML model can estimate  $T_L$  in both clear-sky and partially cloudy days, the derived  $T_L$  from the clear-sky instants in the partially cloudy days as well as corresponding meteorological variables can be included in the dataset for model development and testing. Which in turn can provide more data for ML model training and could potentially improve the model accuracy.

### 3.3. Estimations of daily turbidity and 1-minute averaged DNICs

The same method is applied to estimate 1-min averaged DNICs using the improved  $T_L$  estimations. Since PVLIB uses the same  $T_L$  value for calculating GHICs and DNICs, we use the GHICs-estimated  $T_L$  to estimate DNICs, as shown in Fig. 11. Both the recalculations and estimations have better overall performance than PVLIB, the RMSE is reduced from  $76.40 \text{ W m}^{-2}$  to  $47.16 \text{ W m}^{-2}$  and  $50.77 \text{ W m}^{-2}$ ,



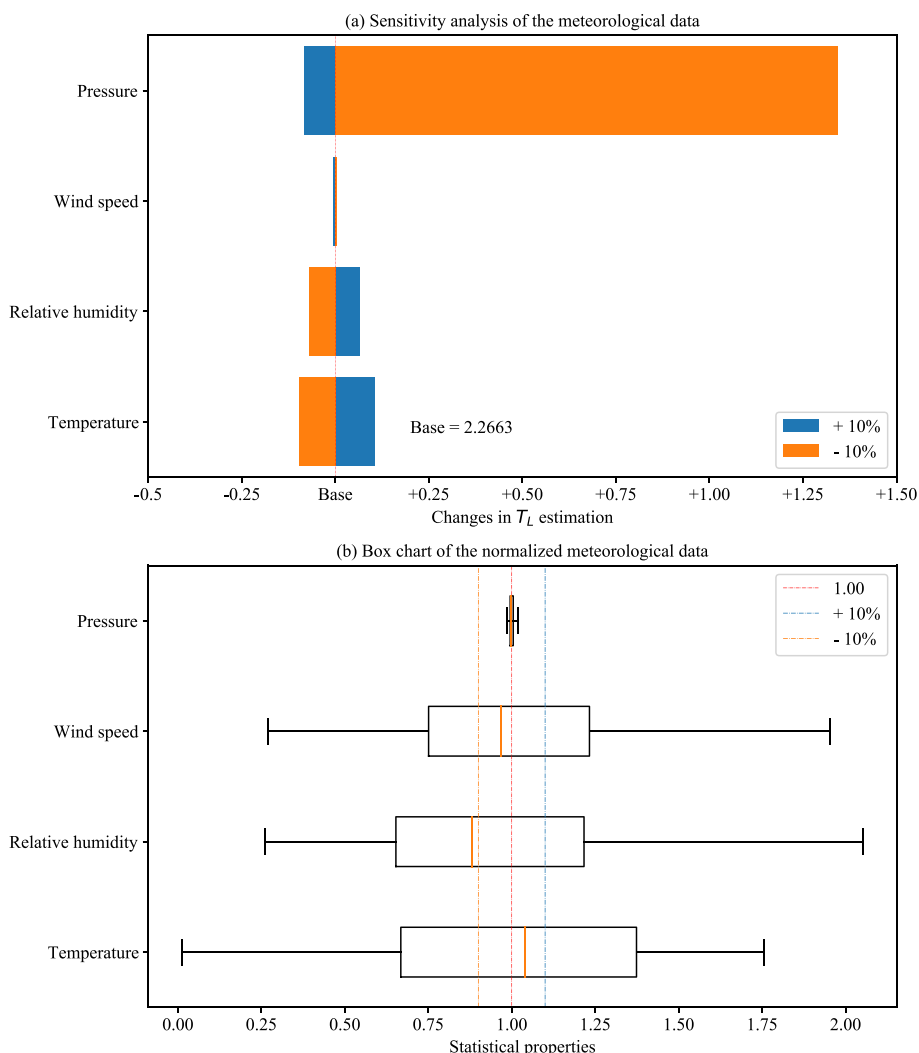


Fig. 8. Sensitivity analysis and statistical properties of the meteorological inputs for the MLP model. (a) Sensitivity analysis based on the changes of a sole parameter, where the base is the mean value. (b) Box chart of the normalized meteorological measurements.

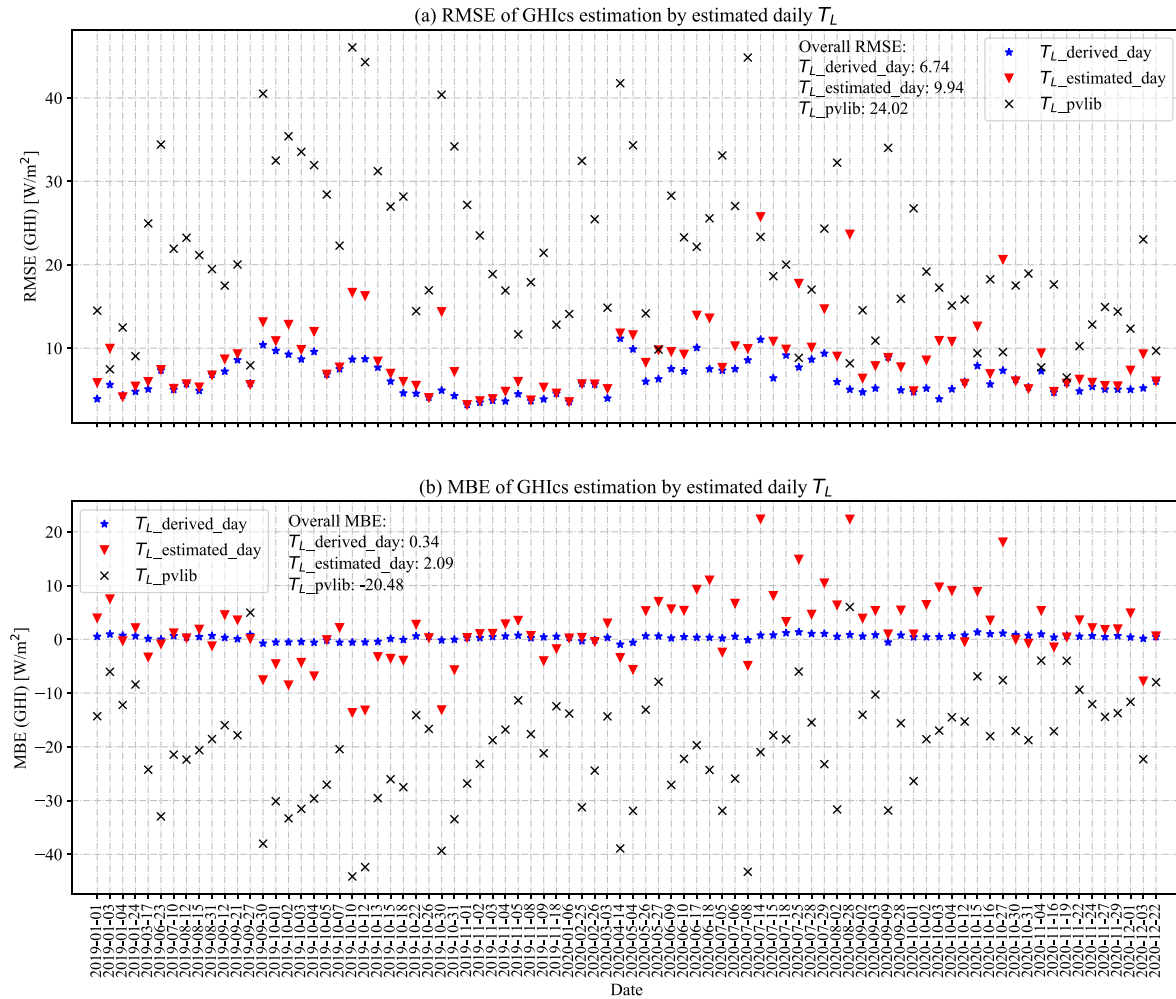
respectively. The MBE is decreased from  $-62.45 \text{ W m}^{-2}$  to  $32.34 \text{ W m}^{-2}$  for recalculations, and to  $39.93 \text{ W m}^{-2}$  for estimations. However, the error reduction is not as effective as GHICs estimation, as it is noticed that the derived  $T_L$  from GHICs could not always lead to better DNICs estimations than PVLIB. Consequently, the estimated  $T_L$  could potentially lead to large errors by accumulating uncertainties in  $T_L$  estimation, as shown in Fig. 11.

To further improve the accuracy of DNICs estimations, we derive  $T_L$  from DNICs and develop separate ML models for  $T_L$  estimation following the similar strategy as described in Section 2.  $T_L$  is derived using Eqs. (2) and (3) from measured clear-sky DNI. A comparison among different  $T_L$  modelling methods for DNICs estimation is shown in Table 3. All the improved  $T_L$  factors for DNICs recalculations and estimations have superior results than default PVLIB. The 5-min averaged  $T_L$  gives the lowest RMSE of  $5.74 \text{ W m}^{-2}$  and a MBE of  $-1.36 \text{ W m}^{-2}$  for recalculating DNICs, while daily mean  $T_L$  generates a RMSE of  $18.93 \text{ W m}^{-2}$  and a MBE of  $1.75 \text{ W m}^{-2}$ . However, the developed MLP models for daily, hourly and 5-min  $T_L$  estimation yield comparable results for estimating DNICs, which means averaging  $T_L$  on smaller time basis has limited potential to improve the DNICs estimation accuracy. Since using daily mean  $T_L$  can generate comparable DNICs estimations with less complexity, a

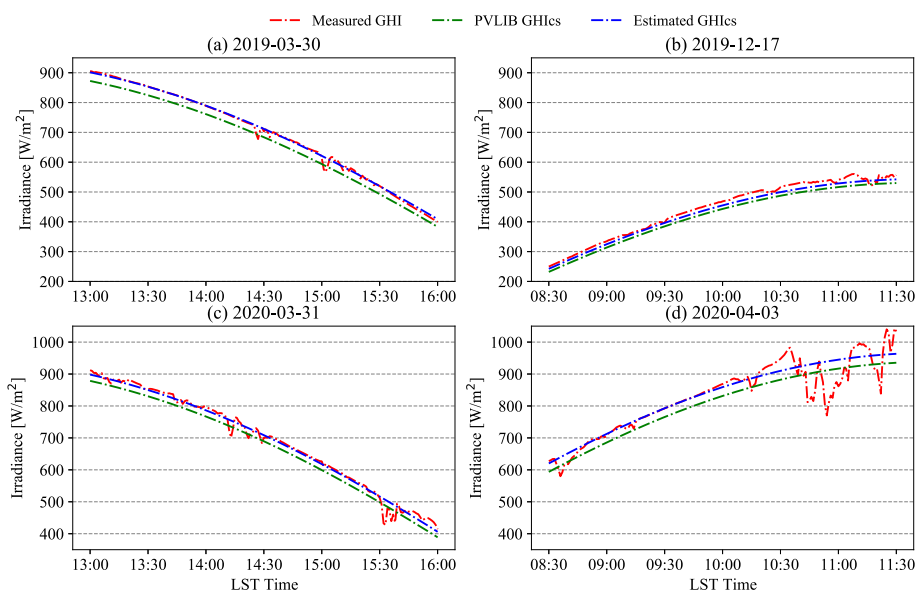
detailed comparison of DNICs estimation by using estimated daily  $T_L$  and default PVLIB  $T_L$  is shown in Fig. 12.

From the perspective of atmospheric radiative transfer, DNI is comparatively more sensitive than GHI to the variations of atmospheric constituents and cloud dynamics, as GHI is the sum of DHI and the horizontal projection of DNI ( $GHI = DNI \cdot \cos(\theta) + DHI$ , where  $\theta$  is the solar zenith angle). The rapidly changing DNICs in the solar morning and evening also makes the DNICs estimation more challenging than GHICs. As demonstrated by our results, the default PVLIB  $T_L$  yields a RMSE of  $76.40 \text{ W m}^{-2}$  and a MBE of  $-62.45 \text{ W m}^{-2}$  for DNICs estimation, which is about three times of the RMSE ( $24.02 \text{ W m}^{-2}$ ) and MBE ( $-20.48 \text{ W m}^{-2}$ ) for estimating GHICs.

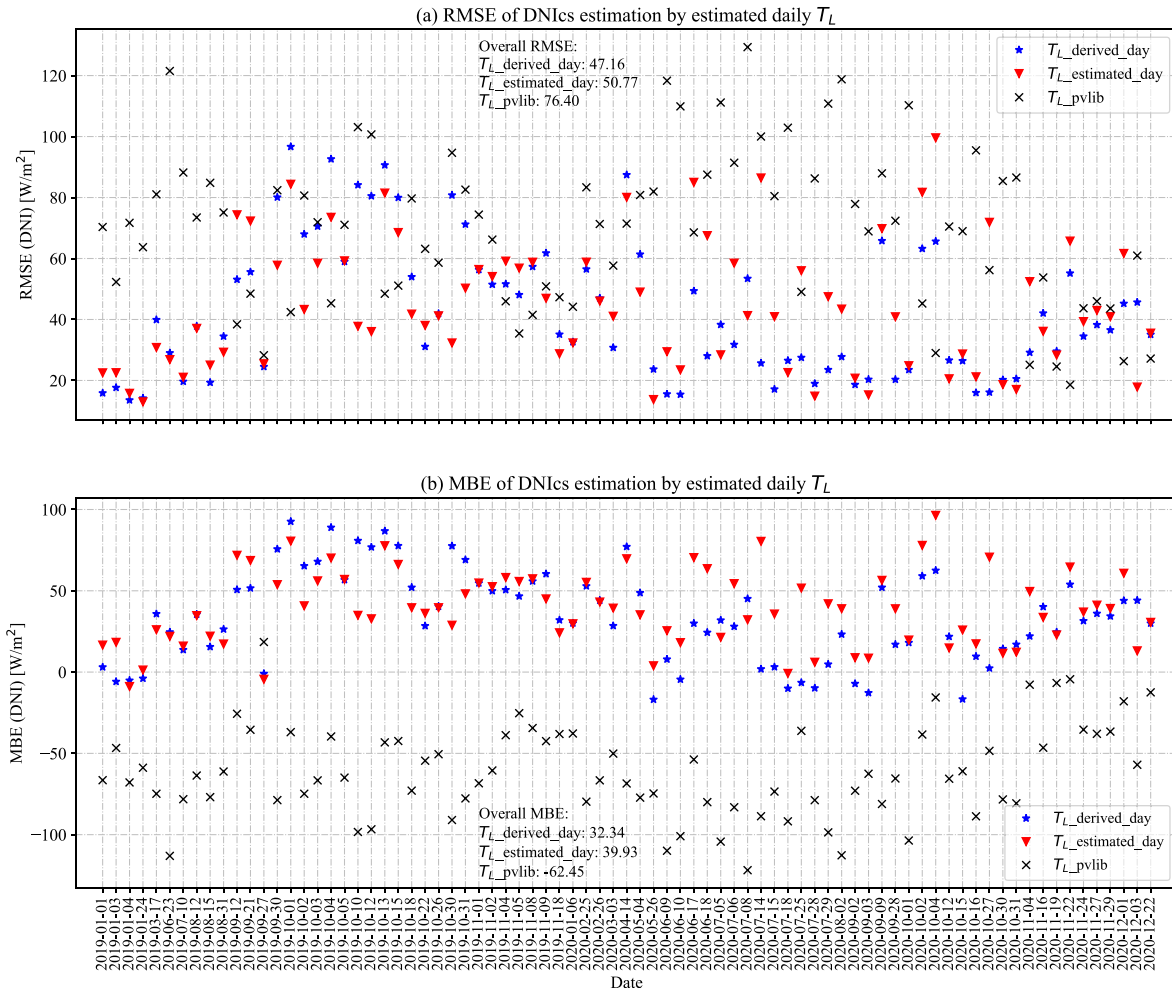
Compared with GHICs estimation from derived and estimated  $T_L$  factors, DNICs estimation generally has comparatively larger errors of RMSE and MBE (see Tables 1 and 3). Using the 5-min averaged  $T_L$  factor almost produce a “perfect” GHICs recalculation with the RMSE of  $0.55 \text{ W m}^{-2}$  and MBE of  $0.01 \text{ W m}^{-2}$ , while the RMSE is  $5.74 \text{ W m}^{-2}$  and MBE is  $-1.36 \text{ W m}^{-2}$  for DNICs recalculation. When it comes to estimation, the ML model (MLP is chosen) estimated daily  $T_L$  for DNICs estimation has a RMSE of  $29.96 \text{ W m}^{-2}$ , which is nearly three times of the RMSE ( $9.94 \text{ W m}^{-2}$ ) of estimating GHICs.



**Fig. 9.** The comparison of GHICs estimation based on derived and estimated daily  $T_L$  factors. (a) Daily RMSE of GHICs estimation. (b) Daily MBE of GHICs estimation. Generally, the estimated  $T_L$  performs better than the default PVLIB  $T_L$  factor.



**Fig. 10.** Examples of GHI in partially cloudy days during (a) 2019-03-30 (b) 2019-12-27 (c) 2020-03-31 (d) 2020-04-03. The GHICs calculated from the estimated  $T_L$  shows a higher accuracy than PVLIB when compared with measured GHI in the clear-sky instants.



**Fig. 11.** The RMSE and MBE of DNics estimation using the GHics-based derived and estimated  $T_L$ . Both recalculations and estimations have lower overall RMSE and MBE than PVLIB but with some exceptions.

**Table 3**

Comparison of DNics recalculations and estimations using derived and estimated  $T_L$  for clear-sky days in the year of 2019 and 2020. Although 5-minute averaged  $T_L$  has the lowest RMSE for DNics recalculations, the 1-min averaged DNics estimations based on daily, hourly and 5-min averaged  $T_L$  show little difference.

$T_L$	DNics recalculations <sup>a</sup>		DNics estimations <sup>b</sup>	
	RMSE [ $W m^{-2}$ ]	MBE [ $W m^{-2}$ ]	RMSE [ $W m^{-2}$ ]	MBE [ $W m^{-2}$ ]
Daily mean	18.93	-1.75	29.96	2.68
Hourly mean	8.96	-1.46	30.75	-0.04
5-min mean	5.74	-1.36	31.96	-1.24
PVLIB <sup>c</sup>	76.40	-62.45	76.40	-62.45

<sup>a</sup> DNics recalculations are based on the averaged  $T_L$  factors derived from DNics.  
<sup>b</sup> DNics estimations are based on the estimated  $T_L$  factors from the ML (MLP) model developed from the derived  $T_L$ .  
<sup>c</sup> PVLIB uses the daily interpolated  $T_L$  based on the monthly climatological  $T_L$  map.

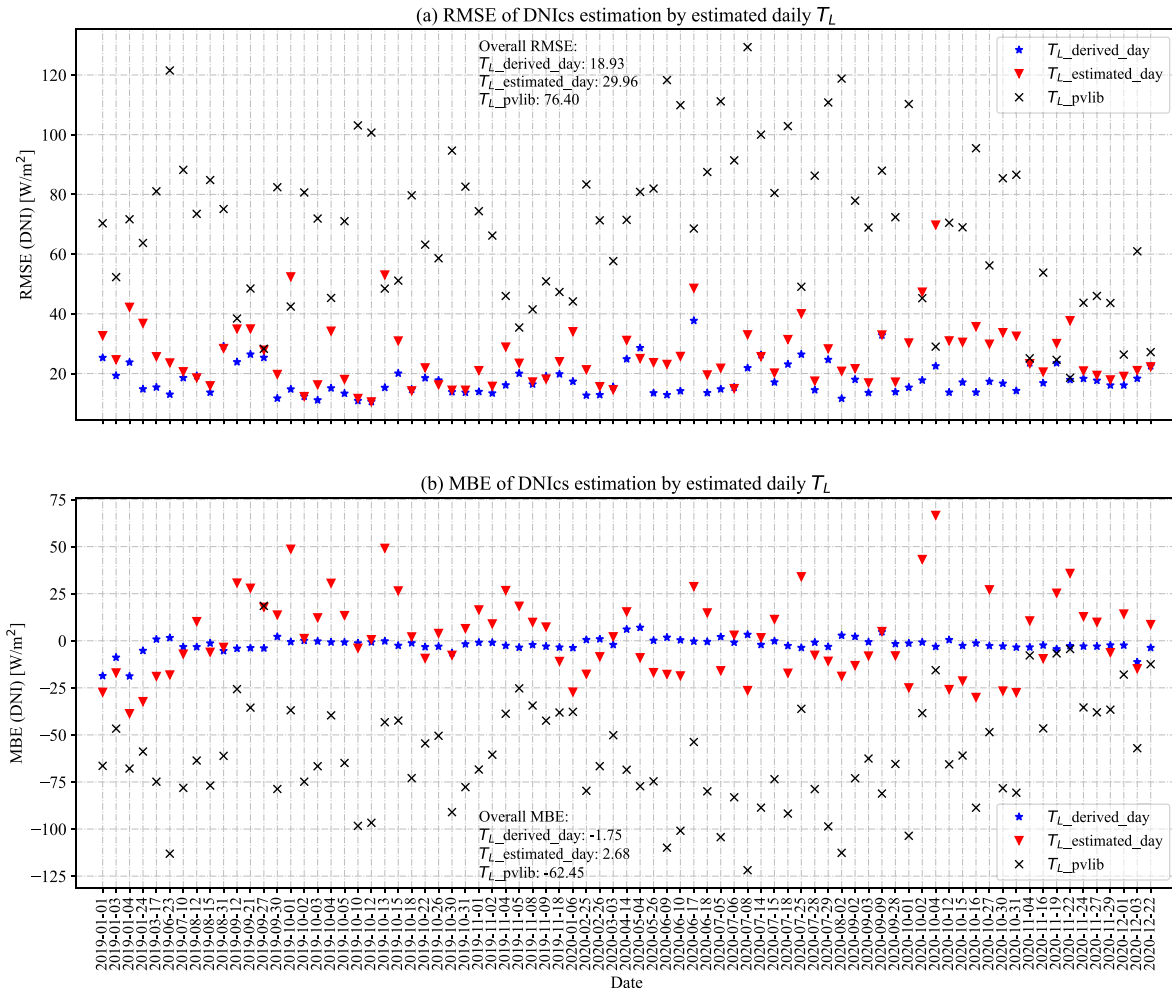
For partially cloudy days, the proposed DNics estimating method also outperforms PVLIB, but the degree of error reduction is smaller than those of GHics estimation, as demonstrated by Fig. 13. In sum, DNics estimation is more challenging than GHics and often has larger discrepancies, the applications that rely heavily on accurate DNics estimation is recommended to adopt the methods to improved DNics estimation (such as the one presented in this work).

### 3.4. Generic applicability of the proposed method

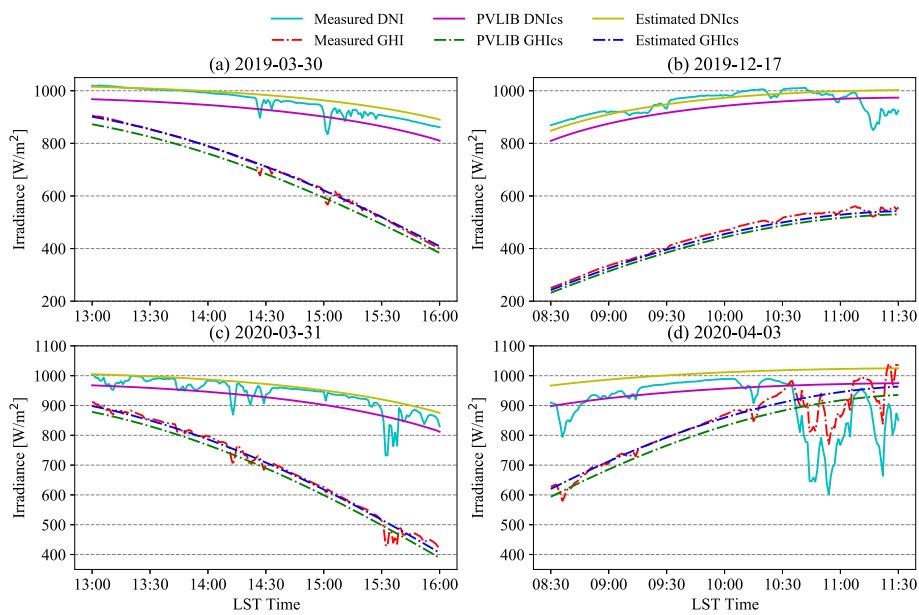
Here we apply the proposed methodology at other SURFRAD stations with limited occurrences of clear-sky days to demonstrate the generic applicability of the proposed method. The results of 1-min averaged GHics estimation for all the SURFRAD stations using the  $T_L$  estimation models developed using both clear-sky and partially clear days are presented in Table 4. Compared with the default PVLIB calculations, the proposed method generally produces better GHics estimations for all SURFRAD stations.

## 4. Conclusions

In this work, we present a new method to estimate turbidity factor  $T_L$  using common meteorological data by ML algorithms. The model inputs are: the default PVLIB  $T_L$ , ambient air temperature, relative humidity (and its logarithm), wind speed, atmospheric pressure, day of year (DOY), and estimated precipitable water. The model output is estimated  $T_L$ , which has the same temporal resolution as the input parameters. The training target of the ML algorithms is the  $T_L$  derived from measured clear-sky GHI or DNI. When tested using data from Desert Rock, Nevada, the new method successfully captures both the short-term and the long-term temporal variations of  $T_L$  by inferring from the local meteorological



**Fig. 12.** The comparison of DNICs estimation based on derived and estimated daily  $T_L$  factors. (a) RMSE of DNICs estimation. (b) MBE of DNICs estimation. Generally, the estimated  $T_L$  performs better than the default PVLIB  $T_L$  factor in terms of RMSE and MBE.



**Fig. 13.** DNI and GHI time series in partially cloudy days during (a) 2019-03-30 (b) 2019-12-17 (c) 2020-03-31 (d) 2020-04-03. Compared with DNICs estimation in partially cloudy days, GHICs estimated from improved  $T_L$  factor has higher accuracy in the clear-sky instants.

**Table 4**

Results of 1-min averaged GHIs estimations using estimated  $T_L$  for clear-sky and partially clear days in 2019 for all the SURFRAD stations. PVLIB results from default  $T_L$  are presented in brackets for reference.

Stations	Clear-sky days		Partially clear days		Clear-sky and partially clear days	
	nRMSE [%]	nMBE [%]	nRMSE [%]	nMBE [%]	nRMSE [%]	nMBE [%]
BON	3.81 (10.68)	−1.90 (−9.08)	3.16 (9.24)	−0.96 (−7.80)	3.38 (9.16)	−0.37 (−7.58)
DRA	1.42 (4.09)	−0.20 (−3.33)	1.48 (3.79)	−0.17 (−3.02)	1.52 (3.80)	−0.13 (−2.99)
FPK	2.90 (7.03)	−1.91 (−4.83)	2.53 (5.32)	−0.78 (−2.97)	2.62 (5.09)	−0.48 (−2.59)
GWN	3.10 (8.52)	−1.19 (−7.52)	3.12 (7.06)	−0.64 (−5.27)	3.24 (7.00)	−0.44 (−4.97)
PSU	1.73 (8.47)	−0.28 (−7.81)	2.29 (7.01)	−0.25 (−6.07)	2.62 (7.11)	−0.35 (−6.04)
SXF	1.69 (7.24)	−0.09 (−6.66)	2.85 (6.13)	−0.03 (−5.01)	3.05 (6.29)	−0.16 (−5.14)
TBL	2.50 (2.73)	1.15 (−0.75)	2.37 (2.45)	1.40 (−0.47)	2.63 (2.66)	1.39 (−0.24)

measurements, thus leading to substantial accuracy improvement in estimating clear-sky irradiance. The major findings and recommendations are:

- We perform  $T_L$  estimation on the averaging basis of a day, an hour and every 5-min. Although 5-min averaged  $T_L$  can better represents its temporal variation, using daily or hourly averaged  $T_L$  to estimate GHIs or DNIs has no significant reduction in accuracy. Therefore, we recommend using the improved daily-averaging  $T_L$  (with less complexity and less computational resource requirement) for GHIs or DNIs estimations.
- Although the default Ineichen-Perez clear-sky model uses the same turbidity factor for GHIs and DNIs estimations, we found that using the same values would deteriorate DNIs estimation. Therefore, we recommend using two separately trained ML models to generate different  $T_L$  values, one for GHIs estimation and one for DNIs estimation.
- During clear days, when compared with the default PVLIB  $T_L$ , the RMSE of GHIs estimation based on the improved daily  $T_L$  decreased from  $24.02 \text{ W m}^{-2}$  to  $9.94 \text{ W m}^{-2}$ , a 58.6% reduction of error. The RMSE of DNIs estimation is reduced from  $76.40 \text{ W m}^{-2}$  to  $29.96 \text{ W m}^{-2}$ , a 60.8% reduction of error. The default PVLIB generally underestimates the GHIs and DNIs with an MBE of  $-20.48 \text{ W m}^{-2}$  and  $-62.45 \text{ W m}^{-2}$ , respectively. The bias are corrected when using the improved daily  $T_L$ , yielding an MBE of  $2.09 \text{ W m}^{-2}$  for estimating GHIs, and  $2.68 \text{ W m}^{-2}$  for estimating DNIs, respectively.
- The daily  $T_L$  estimation method is also tested in partially cloudy days (with partial clear periods and partial cloudy periods). It is observed that the corresponding GHIs and DNIs estimations show better agreement with clear-sky irradiance measurements during cloudless time instances, when compared with default PVLIB results. The results indicate that the presence of clouds does not significantly change local air temperature and relative humidity, as well as water vapor and aerosol concentrations. Furthermore, the results demonstrate the potential of the proposed method in assisting solar irradiance modelling and forecasting in partially cloudy conditions, especially for cloud identification applications.

In sum, our proposed method offers a simpler way for  $T_L$  estimation without priori knowledge of aerosol and water vapor content in the atmosphere. The estimated  $T_L$  can substantially improve the accuracy of clear-sky GHI and DNI estimations when used in an empirical clear-sky model. Our results also imply that local meteorological data such as air temperature and relative humidity can represent column water vapor and aerosol concentrations with high accuracy during both clear and partially cloudy days. Solar resourcing and forecasting applications are expected to be

improved when the proposed method is used to estimate clear-sky irradiance with higher accuracy.

### CRediT authorship contribution statement

**Shanlin Chen:** Methodology, Software, Investigation, Writing – original draft. **Mengying Li:** Conceptualization, Resources, Supervision, Writing – review & editing.

### Declaration of competing interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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